Low-latitude HF Doppler short-period oscillations associated with storm sudden commencements

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HF Doppler measurements of F-region vertical plasma drifts (Vz) at a low-latitude station, Visakhapatnam (dip-17°N) in association with two storm sudden commencements (SC*/SC) on 21 February 1994 (during daytime 0902 h UT) and 28 May 1994 (during night time 1357 h UT) are analysed. Measurements of Vz at SCs are compared with the H-field variations of high latitudes down to dip equatorial latitudes. It is found that there is an abrupt enhancement in Vz (or in Doppler velocity $V_{\rm D}$) in coincidence with the main impulse of SC*/SC in both the events. In the case of the first event, upward enhancement in Vz is unusual and is nearly 2 to 3 times that of the usual velocities. Also, in the same event there is a downward trend in Vz prior to this enhancement, after which short-period oscillations (60–90 s) in Vz which are coherent with the magnetic fluctuations at several high latitudes down to equatorial latitudes are observed. The observed changes in Vz (V_D) at SCs are discussed in the light of the effects of magentosphere-generated electric fields promptly penetrated to these low latitudes by means of transmission of electromagnetic waves through earth-ionosphere waveguide and direct transmission of compressional hydromagnetic waves.

Keywords: Doppler velocity, HF Doppler measurements, oscillations, plasma drift, storm sudden commencement.

MAGNETOSPHERIC disturbances such as storm sudden commencements (SSCs) or simply sudden commencements (SCs) are attributed to the interaction of sudden changes in the solar wind dynamic pressure with the earth's magnetosphere. Hence these disturbances provide an opportunity to study the response of the magnetosphere to solar wind variations.

Among a variety of SCs, a class called as SC* is characterized by a negative impulse in the H-field component called preliminary impulse (PI) preceding the main impulse (MI) 1 . The PI is observable at high and dip equatorial latitudes but may not appear significantly at low latitudes. The ionospheric current system for the PI was first identified by Araki $et\ al.^2$. It was recognized that the PI and MI

correspond to two different driving mechanisms. After a thorough study of SSCs, Araki³ proposed a model to interpret the observed features. According to this model, the components of SC disturbance field are given by $D_{SC} = DP_{PRI} + DP_{MI} + DL_{MI}$, where PRI part of SC and a part of MI are due to polar electric fields and the remaining part of MI is due to compressional hydromagnetic waves. Thus the disturbance of SC may manifest in the *H*-field in the atmosphere as superposition of DP part over DL part.

In addition to the above features of PI and MI, sometimes magnetic disturbances may manifest as short-period fluctuations at the onset of SCs. These fluctuations are called micropulsations and are of two types: pulsations continuous (Pc) and pulsations irregular (Pi). Recently, electric-field fluctuations in coherence with magnetic fluctuations (Pc5, Pi-2) have been observed during magnetospheric disturbances such as DP-2 events and substorms^{4,5}. Doppler oscillations associated with short period (Pc3-4) pulsations at the times of SCs are rarely reported. Low-latitude response during these disturbances may provide additional information in studies of interaction of the magnetosphere with the solar wind dynamical changes. In the present study it has been observed that in one of the SC events, quasi-periodic oscillations are found at the onset of SC. This event provided an opportunity to study the response of low-latitude ionosphere to these magnetic fluctuations, in addition to the general SC features.

HF Doppler technique is a sensitive tool in detecting ionosheric electric field oscillations of several scale sizes during mangnetospheric disturbances such as SC, SI, DP-2 events, etc. $^{6-11}$. In the present investigation HF Doppler measurements of vertical plasma drifts (Vz) are used to study the response of the low-latitude ionosphere during two SC events.

We have recorded phase path of *F*-region reflections at 5.5 MHz during 1993–95 using a coherent HF Doppler radar at the low-latitude station, Visakhapatnam (dip 20°N). Scrutiny of data showed the availability of phase-path measurements during two SC events on 21 February and on 28 May 1994.

The HF Doppler system consists of pulsed HF transmitter, phase-coherent receiver, frequency synthesizer, multichannel strip-chart recorder, and transmitting and receiving antennas. The transmitter pulsed RF, and the receiver local oscillator frequencies as well as the timing and gate pulses were all derived from a 10 MHz crystal oscillator whose stability is better than one part in 10^7 . The system was operated at a frequency of 5.5 MHz with a pulse width of $100 \,\mu s$ and a pulse repetition frequency (PRF) of $50 \, Hz$, to probe the bottom side F-region. The receiver detects the received signal and provides sampled outputs of in phase ($I = A\cos\varphi$) and quadrature ($Q = A\sin\varphi$) signals, where A and φ are the instantaneous amplitude and phase of the received signal. The I and Q signals were recorded on a strip-chart. From these, the time variations of A and

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 φ were scaled. The rate of change of φ gives the Doppler frequency (Δf) of the receiver signal and the line-of-sight plasma drift velocity Vz. The chart recorder was run at a speed of 10 cm/min, which provides measurements of Doppler shifts in the range of 0–2.5 Hz, corresponding to vertical plasma drift velocities in the range of 0–70 m/s. In the present study, the rate of φ was scaled at 12 s intervals and from these values, Doppler velocities (V_D) and vertical plasma drifts ($Vz = V_D/2$) were calculated. (From the group height measurements noted in the log book, it is found that in both the SC events the height of the reflecting layer is well above 300 km and hence chemical losses may be neglected). These V_z (or V_D) values are compared with the H-field variations in the magnetograms of a nearby station, Hyderabad (geomag. lat. 7.6°N) and from World Geomangnetic Data Centers (courtesy of WDC-C2, and 210 Magnetic Meridian Chain).

Scrutiny of solar geophysical data revealed that the SC on 21 February 1994 is reported as SC* and remarkable by 16 observatories, whereas the SC on 28 May 1994 was reported as simple SC by ten observatories. *Vz* and *H*-field variations during these two events are presented.

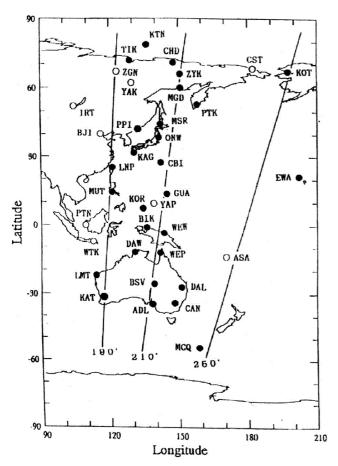
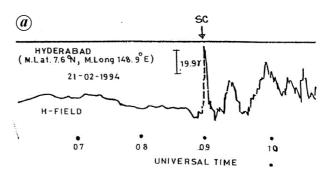


Figure 1. Map showing locations of a chain of stations covering $\pm\,60^\circ$ geographic latitudes around the dip equator and around 190, 210 and 250° geographic longitudes.

For the SC event on 21 February 1994, a chain of magnetic observatories lying nearly around the same longitude (210 MM network stations) is shown in Figure 1. Stations covering high latitudes down to equatorial on both the hemispheres ($\pm 60^{\circ}$ geographic latitudes are also shown). *H*-field variations recorded at a nearby low-latitude station, Hyderabad (M. lat. 7.6°N, M. long 148.9°E) are shown in Figure 2 a. An abrupt increase in the *H*-component, nearly 45 nT at SC (indicated by arrow) around 0900 h UT is also shown and then the field is found to suddenly decrease. The exact duration of rise and fall is not identifiable, as the magnetograms are normal-run magnetograms. The magnetogram also indicates quasi long period (30–40 min) oscillations in the *H*-component after SC, which continued up to 12 h UT (1730 h IST).



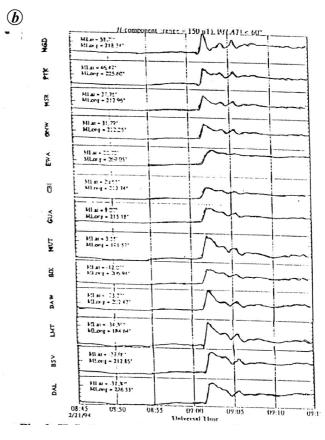


Figure 2. *H*-field variations associated with SC on 21 February 1994 showing SC onset at 0901 h UT (a) at Hyderabad and (b) at several stations covering \pm 60° geographic latitudes around the dip equator.

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Station	Abbreviation	Geographic coordinates		Geomagnetic coordinates	
		Latitude	Longitude	Latitude	Longitude
Magadan	MGD	59.97	150.86	53.56	218.66
St. Paratunka	PTK	52.94	158.25	46.34	225.91
Moshiri	MSR	44.37	142.27	37.61	213.23
Onagawa	ONW	38.43	141.47	31.65	212.51
Ewa Beach	EWA	21.32	202.00	22.67	269.36
Chichijima	CBI	27.15	142.30	20.59	213.00
Guam	GUA	13.58	144.87	4.57	214.76
Muntinlupa	MUT	14.37	121.02	3.58	191.57
Biak	BIK	-1.08	136.05	-12.18	207.30
Darwin	DAW	-12.40	130.90	-23.13	202.68
Learmonth	LMT	-22.22	114.10	-34.15	185.02
Birdsville	BSV	-25.54	139.21	-36.58	212.96
Dalby	DAL	-27.18	151.20	-37.09	226.80

Table 1. Geographic and geomagnetic coordinates of observation sites of magnetograms shown in

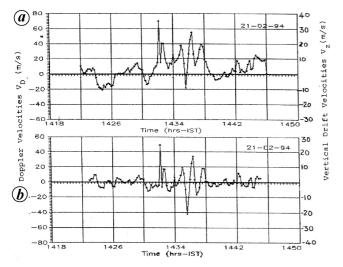


Figure 3. Variations of Doppler velocities of F-region reflections on 5.5 MHz associated with SC onset on 21 February 1994 at Visakhapatnam observed in (a) vertical drifts (Vz) and (b) short-period fluctuations in Vz.

Magnetograms of a few stations of the 210 MM network chain (1 s plots of *H*-field) are shown in Figure 2 b. Geographic and geomagnetic coordinates of these stations are given in Table 1. In all these magnetograms there is a sudden increase in the *H*-field indicating SC and thereafter, quasi short-period fluctuations continue for a few minutes (5 min). Thereafter, the H-component decreases to pre-SC level. Fluctuations in all the magnetograms indicate that these are of global nature. Also, in the magnetograms of high-latitudes stations [MGD, PIK, LMT, BSV, DAL], there is a small negative impulse (PI) before the main impulse (MI). This small-amplitude PI is not visible in lowlatitude magnetograms. Solar geophysical data revealed that 16 observatories reported this SC as SC* and are remarkable. This confirms the features of the event on 21 February 1994 as SC*.

Vz and Doppler velocities on this storm day are shown in Figure 3 a. The Doppler velocity is abruptly enhanced at 0902 h UT (1432 h IST), coincident with the main impulse of the SC. The upward enhancement is from -18 to 65 m/s. This velocity is unusual and is nearly 2 to 3 times that of usual velocities during daytime. After this enhancement, Doppler velocity is suddenly decreased and again increased in the upward direction. Thus, oscillations in Vz (V_D) are continued up to 1438 h IST (0908 h UT). These oscillations are extracted by passing Doppler velocities through a running mean filter (180 s) and the residues are shown in Figure 3 b. Periodicities of the oscillations which are in the range of 60 to 90 are also shown. Magnetic fluctuations observed in all the magnetograms of the network station shown in Figure 2 b and the short-period Doppler oscillations shown in Figure 3 b occurred in coincidence and continued for about 5 min. This indicates that these Doppler oscillations are the effects of global fields generated and/or propagated to these latitudes at the time of SC. But it is found that the Doppler oscillations are not in phase with magnetic fluctuations of any of the stations shown in Figure 2 b.

For the SC event on 28 May 1994 also the 210 MM network record is used. The record on the storm day is shown in Figure 4. It can be seen from Figure 4 that in most of the stations there is an abrupt increase the in *H*-field at around 1400 h UT (1930 h IST), but of small amplitude. Later, the field decreased gradually in a few tens of minutes and then further to below normal levels. In the solar geophysical data, it is found that more than ten observatories reported this SC and the onset at 1357 h UT (1927 h IST). Thus the SC in this event is simple SC having the characteristic of MI only.

The *H*-component magnetogram collected from Hyderabad on this storm day is shown in Figure 5 *a*. In this mangnetogram the SC onset is at around 1400 h UT (1930 h IST) and is of small amplitude (nearly 13 nT).

The SC is indicated with an arrow. After this rise, the H-component gradually decreased in a few minutes and crossed its normal value at around 1430 h UT. The Doppler velocities (V_D) on this day during 1352–1400 h UT are shown in Figure 5 b. It can be seen from Figure 5 b but there is a sudden increase in V_D in coincidence with the onset of SC. Though the magnitude is small, there is a sudden increase in V_D from -2 m/s to +2 m/s at 1927 h IST. This sudden increase is coincident with the MI of the SC.

In the present study it is confirmed that there is sudden enhancement in Vz (or V_D) in coincidence with the main impulse (MI) of the SC (0902 h UT). Also, prior to this sudden increase, there is a downward trend in Vz, which may be at the time of PI. Thus the downward trend in Vzbefore MI, may be attributed to the effect of PI, though this downward drift is not significant. At these low-latitude stations, vertical plasma drifts are mainly due to effects of electric fields and/or neutral wind fields. It has been confirmed by several authors that changes in the H-field at MI and PRI are due to the effects of magnetosphericgenerated electric fields^{3,6,7}. As there is no time delay between changes in Vz and changes in magnetic field at global stations, the observed changes in Vz at MI and PRI can be attributed to the effects of magnetospheric-generated electric fields.

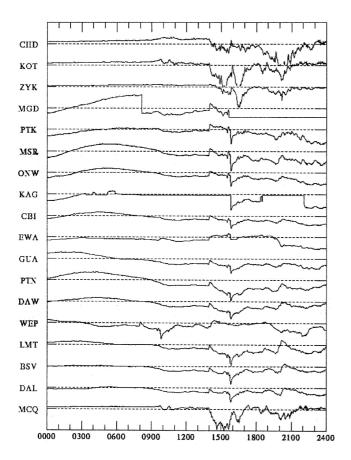


Figure 4. *H*-field variations showing SC on 28 May 1994 at a chain of stations covering $\pm 60^{\circ}$ geographic latitudes around the dip equator.

Kikuchi⁶ provided the evidence for the prompt penetration of dawn-to-dusk eastward electric field to low-latitudes at the time of MI and dusk-to-dawn westward electric field at the time of PRI of SC* event. Sastri *et al.*⁷ using HF Doppler measurements at equatorial station, provided evidence for penetration of eastward electric field in response to PI of SC* and westward electric field in response to MI.

In the present study, the observed features of enhancement in the upward drift in Vz coincident with MI and a small downward drift just before this enhancement are in agreement with the results of Kikuchi. As the amplitude of the magnetospheric-generated electric field decreases when it penetrates down to low latitudes, the effect of the electric field responsible for PRI may not be significant at this low-latitude station. Hence the small downward drift before MI may be attributed to the effect of electric fields due to PRI. Thus the upward drift at MI and downward drift at PRI may be due to presence of eastward and westward electric fields. As suggested by Kikuchi⁶, the eastward electric fields at MI might be generated in the polar region and are instantaneously propagated to these low latitudes by electromagnetic waves through earthionosphere waveguides and the presence of westward electric field at PRI of SC* might be the induced field due to direct transmission of compressional hydromagnetic waves.

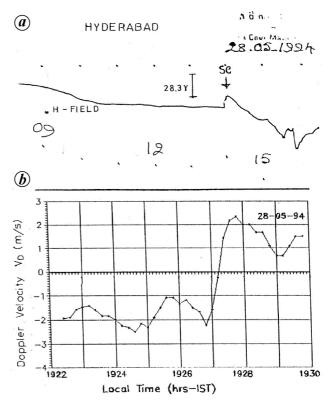


Figure 5. Variations associated with SC onset on 28 May 1994 in (a) horizontal component of magnetic field (H) at Hyderabad and (b) vertical drift velocities (Vz) at Visakhapatnam.

In the case of the first event, the observed short-period (60-90 s) Doppler oscillations coincident with the magnetic fluctuations at observatories located at several longitudes and latitudes around dip equator, may be due to the effect of fields of disturbance of global origin. These magnetic fluctuations called micropulsations, might have occurred at the SC. Matsushita suggested that the initial shock of SC may produce not only hydromagnetic waves responsible for SCs and SIs, but also smaller amplitude and short-period hydromagnetic waves. These waves may propagate simultaneously and oscillate the ionosphere, causing ionospheric currents. From the theory of micropulsations, when the hydromagentic wave moves through the plasmasphere along a flux tube, the varying magnetic field would compress and rarify the intervening plasma. This mechanism may create magnetic fluctuations as well as plasma motions. In a modelling work of Doppler osciallations due to micropulsations, Poole et al. 12 suggested that compressions and rarefractions of the plasma may be the dominant factor for the observed oscillations in the Doppler shifts, when compared to the effects due to refractive index variations. If this were to be considered, as the plasma motions may manifest in EXB drifts, the observed short-period oscillations in V_z (or in V_D) may be because of the effects of plasma motions due to hydromagnetic waves.

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Climatic variability and its impact on cereal productivity in Indian Punjab

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Dynamic crop growth simulation models CERES-Rice and CERES-Wheat for rice and wheat, respectively were used to study the effect of climate change on growth and yield of these crops under non-limiting water and nitrogen availability. Analysis of recent 30 year historical weather data from different locations in the state revealed that the minimum temperatures have decreased or increased (-0.02 to + 0.07 °C/year), maximum temperatures decreased (-0.005 to -0.06°C/ year) and rainfall increased (2.5-16.8 mm/year). Keeping in view the observed trends in climate variability, growth and yield of crops were simulated under plausible synthetic climatic scenarios of changes in temperature and solar radiation. In general, with an increase in temperature above normal, the phenological development in wheat was advanced, but that of rice was not much affected. With an increase in temperature up to 1.0°C the yield of rice and wheat decreased by 3 and 10%, respectively. On the other hand, crop yields decreased with decrease in radiation and vice-versa. The interaction effects of simultaneous increase/decrease in parameters were also simulated. When the maximum temperature decreased by 0.25 to 1.0°C while minimum temperature increased by 1.0 to 3.0°C from normal, the yield in rice and wheat decreased by 0.8 and 3.0%, respectively from normal.

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