



Figure 4. A distinct NNE-SSW trending pressure ridge comprising Quaternary fan sediments. This is the direction of active tectonics observed in the valley (near Dhaniyakot).

strike-slip fault, indicating that the movement is still continuing. Another evidence of active tectonics in the Kosi Valley is well depicted in the form of down fan movement, steep fan scarps, uplifted terraces, pressure ridge (Figure 4) and landslides along the cliff surface of the Quaternary sediments. The faults trending NNE-SSW being the youngest, displace all the structural elements of the area. The NNE-SSW trending faults recorded in the Quaternary sediments indicate that these faults are also active besides longitudinal faults/thrusts.

Thus Quaternary sequences developed in the Kosi Valley around Betalghat have provided ample evidence of active tectonics. The uplifted and truncated fluvial terraces, tectonic landforms and offsetting in the Quaternary sediments along transverse faults indicate that the Kumaun Lesser Himalaya is tectonically active. Although preliminary in nature, the available field evidences as discussed above indicate that the transverse faults have developed on account of oblique convergence of the Indian plate under the Eurasian plate. However, a more definitive inference should await a detailed morphotectonic study of the area.

1. Bagati, T. N. and Thakur, V. C., Quaternary basins of Ladakh and Lahul-Spiti in north-western Himalaya. *Curr. Sci.*, 1993, **64**, 898–903.
2. Valdiya, K. S., Uplift and geomorphic rejuvenation of the Himalaya in the Quaternary period. *Curr. Sci.*, 1993, **64**, 873–885.
3. Yeats, R. S., Nakata, T., Farah, A., Fort, M., Mirza, M. A., Pandey, M. R. and Stein, R. S., The Himalayan Frontal Fault system. *Ann. Tecton.*, 1992, **6**, 85–98.
4. Valdiya, K. S., The Main Boundary Thrust zone of the Himalaya. *Ann. Tecton.*, 1992, **6**, 54–84.
5. Bilham, R. and Gaur, V. K., Geodetic contribution to the study of seismotectonics in India. *Curr. Sci.*, 2000, **79**(9), 1259–1269.
6. Bilham, R., Blume, F., Bendick, R. and Gaur, V. K., Geodetic constraints on the translation and deformation of India: Implication for future great Himalayan earthquake. *Curr. Sci.*, 1998, **74**, 213–229.
7. Wobus, C., Heimsath, A., Whipple, K. and Hodges, K., Active out-of-sequence thrust faulting in the central Nepalese Himalaya. *Nature*, 2005, **434**, 1008–1011.

8. Srivastava, P., Tripathi, J. K., Islam, R. and Jaiswal, M. K., Fashion and phases of late Pleistocene aggradations and incision in the Alaknanda River valley, western Himalaya, India. *Quaternary Res.*, 2008, **70**, 68–80.
9. Juyal, N., Sundriyal, Y. P., Rana, N., Chaudhary, S. and Singhvi, A. K., Late Quaternary aggradation and incision in the monsoon-dominated Alaknanda valley, Central Himalaya, Uttarakhand, India. *J. Quaternary Sci.*, 2010, **25**, 1293–1304.
10. Jayangondaperumal, R., Dubey, A. K. and Sen, K., Structural and magnetic fabric studies of recess structures in the Western Himalaya: Implications for 1905 Kangra earthquake. *J. Geol. Soc. India*, 2010, **75**, 225–238.
11. Kandpal, G. C., Joshi, K. C., Joshi, D. D., Singh, B. K. and Singh, J., Signature of Quaternary tectonics in a part of Dehra Dun valley, Uttaranchal. *J. Geol. Soc. India*, 2006, **67**, 147–150.
12. Devrani, U. and Dubey, A. K., Anisotropy of magnetic susceptibility and petrofabric studies in the Garhwal Synform, Outer Lesser Himalaya: evidence of pop-up klippen. *Island Arc*, 2009, **18**, 428–443.
13. Rautela, P. and Sati, D., Recent crustal adjustments in Dehra Dun valley, western Uttar Pradesh, India. *Curr. Sci.*, 1996, **71**, 776–780.

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Intense deep convective mixing in the southeast Arabian Sea linked to strengthening of the northeast Indian monsoon during the middle Pliocene (3.4 Ma)

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The climate of the Indian Ocean is dominated by monsoon reversals, influencing hydrography and biogeochemistry of the Indian Ocean as well as land vegetation through changes in precipitation. During summer or southwest monsoon season, intense upwelling zones driven by Ekman spiral appear in the western and eastern parts of the Arabian Sea that enhance surface primary production and thus proliferation of distinct fauna and flora. During the winter season, northeast monsoon winds cause deep convective overturning (mixing) that injects nutrients to the surface ocean and increases surface production. As a result, the primary production in the Arabian Sea has bimodal annual distribution. The present study analyses 5.6 Ma

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record of surface-dwelling planktic foraminifera, *Globigerina bulloides*, *Globigerinoides ruber* and *Globigerinoides sacculifer* from Deep Sea Drilling Project Site 219, southeast Arabian Sea to understand changes in the surface ocean as driven by the Indian monsoon coinciding with the northern hemisphere glaciation (NHG). An increase in mixed-layer species at ~3.4 Ma suggests intense deep convective overturning caused by strong NE monsoon winds related to strengthening of NHG. *Globigerina bulloides* shows a high positive relation with *G. ruber* during the past 3.4 Ma and a weak relation in the early Pliocene (5.6–3.4 Ma). The high *G. bulloides* percentages during the past 3.4 Ma could be linked to the injection of nutrients in the top layer by the advecting sub-surface nutrient-rich water.

Keywords: Deep convective mixing, NE monsoon, southeast Arabian Sea.

THE Pliocene is marked by a transition from a 'greenhouse world'¹ during 5–3.2 Ma to a world marked by periodical waxing and waning of ice sheets (since 3.2–2.4 Ma), when the northern hemisphere witnessed major ice volume expansion². During the early Pliocene warmth, the tropics were in a permanent El Niño state with weak Walker circulation and deep thermocline³. The global climates underwent major reorganization during the middle Pliocene (3.2–2.4 Ma), marked by major expansion of the northern hemisphere glaciation (NHG), leading to more abrupt changes in climate and ocean circulation^{2,4}. Intensification of NHG may have significantly cooled deep waters (formed at high latitudes) that subsequently increased deep ocean stratification^{3,5}. Increased ocean stratification caused ventilated thermocline to shoal, allowing cold water to upwell in the tropics and subtropics, thereby enhancing global cooling³.

Intensification of NHG drove significant changes in land vegetation, triggered African aridification and increased Indian monsoon seasonality with more intense winter or northeast monsoon^{6–11}. During cold intervals the wind intensity (including NE monsoon) increases, which leads to intense convective mixing (overturning) at high and low latitudes. Intense NE monsoon winds driven by NHG could have brought significant changes in surface oceanography, and fauna and flora of the northern Indian Ocean. To understand the impact of the changing monsoonal seasonality, in response to NHG intensification, on surface ocean in the southeast Arabian Sea, this study analysed 5.6 million-year-old proxy record of surface-dwelling planktic foraminifera, *Globigerina bulloides*, *Globigerinoides ruber* and *Globigerinoides sacculifer* from Deep Sea Drilling Project (DSDP) Site 219. We selected these planktic foraminifera because of their dominance at Site 219 and their preference to distinct water masses.

Site 219 was drilled during DSDP Leg 23 on the crest of the Laccadive-Chagos Ridge, off the southwest coast

of India, SE Arabian Sea (9°01.75'N, 72°52.67'E; water depth 1764 m; Figure 1). At present, this site lies within an area where surface primary production is higher due to summer or southwest monsoon-induced upwelling and is thus good for the study of monsoon-driven changes in the surface ocean (Figure 2). However, during NE monsoon season, surface production decreases significantly.

We analysed 111 core samples of 10 cm³ volume from a 68 m long sediment sequence from Site 219 spanning the last 5.6 Myr. Samples were soaked in water with half

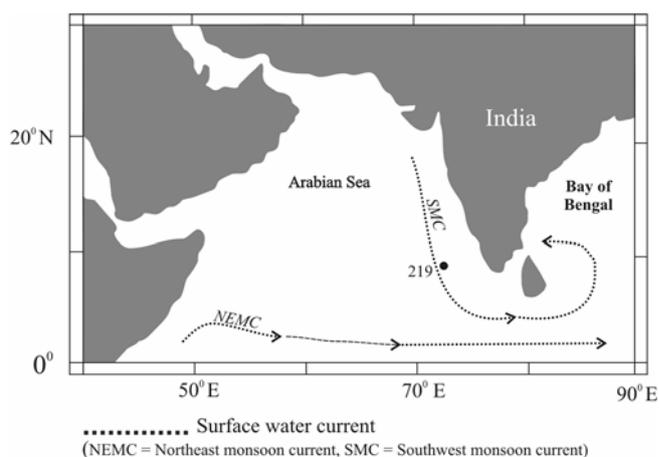


Figure 1. Location map of DSDP Site 219 in the tropical Indian Ocean. The surface ocean currents are from Schott *et al.*⁴¹.

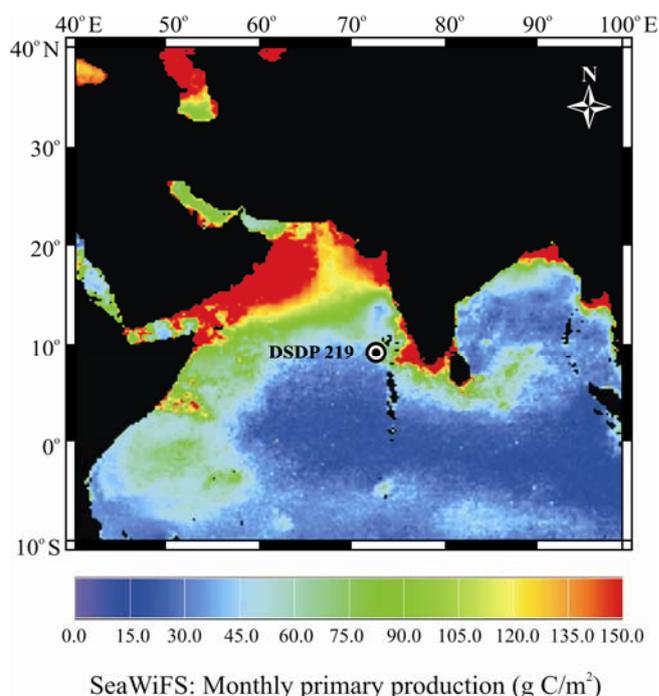


Figure 2. Location of DSDP Site 219 superimposed on the annual primary production map of the Indian Ocean based on SeaWiFS chlorophyll data (averaged over June 1998–August 1998, from http://marine.rutgers.edu/opp/swf/Production/results/all2_swf.html). At present Site 219 lies close to summer monsoon-driven upwelling zone.

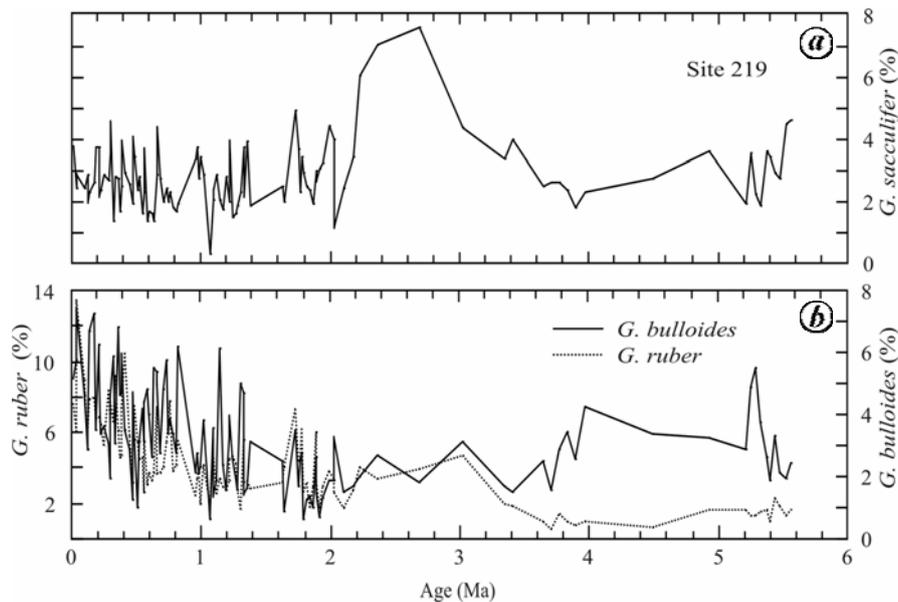


Figure 3. Percentage distribution of planktic foraminifera, *Globigerina bulloides*, *Globigerinoides ruber* and *Globigerinoides sacculifer* during the past 5.6 Ma. *G. bulloides* and *G. ruber* populations deviate out in the early Pliocene (5.6–3.4 Ma) and thereafter show good parallelism contemporaneous with the Northern Hemisphere Glaciation. *G. sacculifer* shows a hump-shaped increase in its population during 3.4–2.2 Ma.

a spoon of baking soda for 8–12 h, and washed over a 63 μm size sieve. The washed samples were dried in an electric oven at $\sim 50^\circ\text{C}$ and transferred into labelled glass vials. Hard sediment samples were soaked in water with 2–3 drops of 2% hydrogen peroxide. We generated census data of planktic foraminifera *G. bulloides*, *G. ruber* and *G. sacculifer* from an aliquot of ~ 300 specimens from $>150 \mu\text{m}$ size fraction. The percentage distribution of each species is shown in Figure 3. The $>150 \mu\text{m}$ size fraction allows us to compare our results with those from the other ocean basins.

Linear correlation between *G. bulloides* and *G. ruber* was carried out to understand the response of the upper ocean in the SE Arabian Sea to changes in monsoonal wind intensities over the past 5.6 Ma (Figure 4). The correlation was carried out for the whole 5.6 Ma interval (Figure 4a), as well as 5.6–3.4 Ma and 3.4–0 Ma intervals (Figure 4b and c) with value of correlation coefficient (R) calculated in each case. The sampling interval is at low resolution during 5.6–3.4 Ma than during 3.4–0 Ma owing to a change in sediment accumulation rate. We could not do much since samples were procured at near equal interval for a different study and were pre-washed.

We applied the age model for Site 219 based on planktic foraminiferal datums suggested in Gupta and Thomas¹². The sediment accumulation rate at Site 219 was high from 5.6–3.4 Ma and moderate from ~ 3.4 –2.4 Ma, with a major increase since ~ 2.4 Ma (ref. 12). The interpolated numerical ages are updated to the Berggren *et al.*¹³ time-scale, with an average time resolution of ~ 50 kiloyear (kyr) per sample.

Globigerina bulloides d'Orbigny is a near-surface (upper 100 m) planktic foraminifer, conventionally known from the transitional and sub-polar water masses¹⁴, but has also been found in significant proportions in tropical and subtropical wind-driven upwelling regions of the northern Indian Ocean¹⁵. This species produces high shell fluxes in monsoon-driven high-productivity regions of the tropical NW Indian Ocean, including the Arabian Sea, and has widely been used in determining SW monsoon wind intensities during the late Quaternary and the Holocene^{15–18}. Sediment trap studies off Somalia indicate highest shell fluxes of *G. bulloides* (21–54%) during summer (SW) monsoon season owing to intense upwelling of cold and nutrient-rich waters^{19,20}. Curry *et al.*²¹ also observed maximum shell fluxes of *G. bulloides* below SW monsoon-induced upwelling zones in the Arabian Sea. During the non-upwelling or intermonsoon periods, however, *G. bulloides* only makes up to 5–12% of the assemblage²⁰.

Globigerinoides ruber is a spinose planktic foraminifer living in the photic zone (top 50 m) of the water column in the tropical and subtropical areas with a sea-surface temperature (SST) of 14–30°C (optimum 21–28°C), and optimum sea-surface salinity (SSS) of 34.5–36.0 psu (ref. 22). It is a symbiont-bearing species (with zooxanthellae), preferring oligotrophic regions with a deep mixed layer²³ and is susceptible to dissolution^{14,24}. *G. ruber* most commonly occupies the warm mixed layer above the thermocline²⁵ and shows maximum abundance in the top 20 m of the mixed layer in the early autumn when thermocline begins to break down²⁶. In the Arabian Sea,

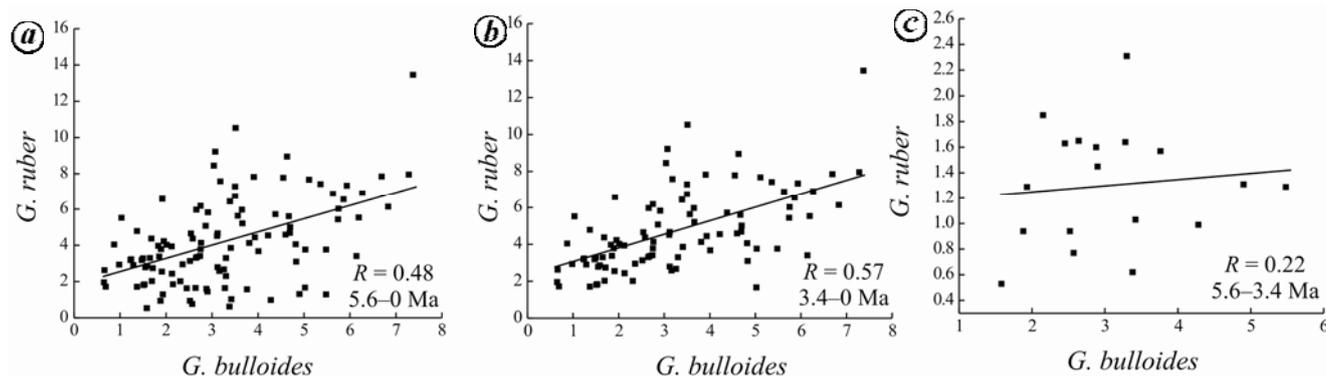


Figure 4. Linear correlation between *G. bulloides* and *G. ruber* over the past 5.6 Ma. The value of correlation coefficient (R) is 0.48 for the whole 5.6 Ma period (a), 0.57 for the interval 3.4 to the Recent (b) and 0.22 for 5.6–3.4 Ma interval (c).

G. ruber proliferates in relatively warm and oligotrophic surface waters during both the upwelling and non-upwelling (both monsoon) seasons owing to its preference to a higher optimum temperature, and the fact that this species may obtain nutrition from its symbionts²⁰. Non-upwelling species produce high or maximum fluxes during the SW monsoon and continue to persist during the non-upwelling periods²⁰.

Globigerinoides sacculifer is also a shallow-dwelling (surface depth habitat top 30–50 m) mixed layer, tropical to subtropical species, tolerating a SST range of ~17–30°C (optimum 27–30°C), SSS optimum of ~34.9 psu and phosphate content of ~0.5 $\mu\text{g l}^{-1}$ (ref. 22). This species prefers to live in warm-water, low-salinity, mixed-layer oligotrophic settings^{27,28}. The species is susceptible to dissolution and prefers low seasonal changes in SST and vertical temperature gradients, and is not well suited to large seasonal salinity changes¹⁴. Duplessy *et al.*²⁹ have suggested that *G. sacculifer* secretes its last chamber below the thermocline. *G. ruber* (white) and *G. sacculifer* mirror regions of highest surface water temperatures and low primary production in the central and southern Arabian Sea³⁰. Oberhänsli *et al.*³¹ observed increased percentages of *G. sacculifer* with increasing oxygen content and slightly lowered salinity in the South Atlantic. It has been observed that when the depth of thermocline is shallow mixed-layer species like *G. sacculifer* decreases²⁸.

Globigerina bulloides and *G. ruber* show a weak positive correlation ($R = 0.22$) from 5.6 to 3.4 Ma, following which the two species show good parallelism ($R = 0.57$; Figure 3). *G. sacculifer*, on the other hand, shows increased percentages from 3.4 to 2.2 Ma, when *G. bulloides* shows a decrease. The value of R between *G. bulloides* and *G. ruber* is 0.48% for the whole 5.6 Ma period (Figure 4). The good correlation between *G. bulloides* and *G. ruber* since 3.4 Ma suggests that the character of the upper ocean in the SE Arabian Sea changed due to a switch in the monsoon wind intensity following strengthening of the NHG.

In the Arabian Sea, where summer monsoon induces upwelling, surface primary production is generally higher

during interglacial times^{16,32,33}. On the other hand, Rostek *et al.*³⁴ observed higher NE monsoon-driven palaeoproductivity in the eastern Arabian Sea caused by deeper mixing during glacial times. Today, NE monsoon winds cause deep convective overturning (mixing) in the Arabian Sea, thereby injecting nutrient-rich subsurface waters into the euphotic zone^{34–37}, which results in a moderate increase of surface temperature. Veldhuis *et al.*³⁸ suggested that wind-induced mixing results in entrainment of nutrient-rich deeper water to the surface, resulting in increased production of phytoplankton and biomass. During the 1995 US Joint Global Ocean Flux Study, Hansell and Peltzer³⁹ observed highest total organic carbon concentrations in the mixed layer during the NE monsoon period, which serves as a source of nutrients to the euphotic zone. Elevated levels of surface chlorophyll and primary productivity are thus associated with deep convective mixing and entrainment during the two monsoon seasons⁴⁰.

Population trend of surface-dwelling planktic foraminifera at Site 219 suggests a major shift in the physical character of the surface ocean in the SE Arabian Sea at ~3.4 Ma, indicating a change in monsoon wind intensities. *G. bulloides* and *G. ruber* show almost no correlation during 5.6–3.4 Ma, with moderate *G. bulloides* and decreased *G. ruber* percentages, indicating increased surface stratification. The two species show good parallelism since ~3.4 Ma, indicating an intense deep convective overturning and entrainment of nutrients in the surface ocean due to stronger NE monsoon winds driven by intense NHG. At present very cold winters produce deep convection in the northern North Atlantic, and in general, the colder the winter, the greater is the overturning.

The fact that *G. ruber* proliferates in both the monsoon seasons whereas *G. bulloides* blooms during summer upwelling (high nutrient level), it is likely that the parallelism between the two species was driven by the availability of nutrients due to deep convective overturning, and not upwelling, during strong NE monsoon season. In a typical upwelling setting, *G. bulloides* population increases to more than 20% and can reach up to ~50% of the total population¹⁷. This suggests that *G. bulloides*

population can also increase during intense winter monsoon that entrains chlorophyll-rich subsurface water to the surface, as it does in the present-day ocean. An increase in *G. sacculifer* during 3.4–2.2 Ma perhaps indicates low surface salinities due to increased NE monsoon precipitation.

During the middle Pliocene, the monsoon entered a regime marked by major increase in NHG superimposed by variations in the earth's orbital parameters¹¹. A change in monsoon seasonality towards intense winter monsoon is observed during 3–2.5 Ma in the eastern Indian Ocean, coinciding with major expansion of the NHG¹¹. It has been observed that during cold intervals the summer monsoon weakens and winter monsoon winds become stronger^{6,10,11}. The monsoon underwent more rapid changes during 3.6–2.6 Ma (ref. 10). Thus there is considerable evidence of linkage between NHG and the development of, as well as fluctuations in, the monsoonal system. Gupta and Thomas¹¹ demonstrated that such weakening of the summer monsoon and initiation of a strong winter monsoon has profoundly affected biota in the eastern equatorial Indian Ocean.

We suggest that monsoon regimes over Site 219 switched between SW and NE monsoons on glacial–interglacial timescale with more influence of the SW monsoon during warm intervals and of the NE monsoon during cold intervals. Since ~3.4 Ma, surface productivity over Site 219 was driven by NE monsoon winds. This study strengthens earlier observations that in the Indian Ocean, monsoon wind regimes changed on glacial–interglacial timescale³². This study also suggests that *G. bulloides*, being an opportunist, responds to the input of cold, nutrient-rich waters, irrespective of the mechanism that brings nutrients to the surface. Although in areas of deep convective overturning the *G. bulloides* population is not as high as in upwelling areas, its population certainly shows a positive correlation with the intensity of the convective overturning.

1. Budyko, M. I., Ronov, A. B. and Yanshin, A. L., *The History of the Earth's Atmosphere*, Gidrometeoizdat, Leningrad, English translation, 1987, Springer Verlag, Berlin, 1985, p. 139.
2. Shackleton, N. J. et al., Oxygen isotope calibration of the onset of ice-rafting and history of glaciation in the North Atlantic region. *Nature*, 1984, **307**, 620–623.
3. Ravelo, A. C., Andreasen, D. H., Lyle, M., Lyle, A. O. and Wara, M. W., Regional climate shifts caused by gradual global cooling in the Pliocene epoch. *Nature*, 2004, **429**, 263–267.
4. Raymo, M. E., The initiation of Northern Hemisphere Glaciation. *Annu. Rev. Earth Planet. Sci.*, 1994, **22**, 353–383.
5. Raymo, M. E., Hodell, D. and Jansen, E., Response of deep ocean circulation to initiation of northern hemisphere circulation (3–2 Ma). *Paleoceanography*, 1992, **7**, 645–672.
6. Fontugne, M. R. and Duplessy, J.-C., Variations of the monsoon regime during the Upper Quaternary: evidence from carbon isotopic record of organic matter in North Indian Ocean sediment cores. *Palaeogeogr., Palaeoclimatol., Palaeoecol.*, 1986, **56**, 69–88.
7. Thompson, R. S., Pliocene environments and climates in the western United States. *Quaternary Sci. Rev.*, 1991, **10**, 115–131.
8. Kukla, G. and An, Z., Loess stratigraphy in central China. *Palaeogeogr., Palaeoclimatol., Palaeoecol.*, 1989, **72**, 203–225.
9. deMenocal, P., Plio-Pleistocene African climate. *Science*, 1995, **270**, 53–59.
10. Zhisheng, A., Kutzbach, J. E., Prell, W. L. and Porter, S. C., Evolution of Asian monsoons and phased uplift of the Himalaya–Tibetan plateau since late Miocene times. *Nature*, 2001, **411**, 62–66.
11. Gupta, A. K. and Thomas, E., Initiation of Northern Hemisphere glaciation and strengthening of the northeast Indian monsoon: Ocean Drilling Program Site 758, eastern equatorial Indian Ocean. *Geology*, 2003, **31**, 47–50.
12. Gupta, A. K. and Thomas, E., Latest Miocene through Pleistocene paleoceanographic evolution of the northwestern Indian Ocean (DSDP Site 219): Global and Regional factors. *Paleoceanography*, 1999, **14**, 62–73.
13. Berggren, W. A., Kent, D. V., Swisher, C. C. and Aubry, M. P., A revised Cenozoic geochronology and chronostratigraphy. In *Geochronology Timescales and Global Stratigraphic Correlation: Framework for and Historical Geology* (eds Berggren, W. A. et al.), Society of Economic Paleontologists and Mineralogists (Society for Sedimentary Geology) Special Publication, 1995, vol. 54, pp. 129–212.
14. Bé, A. W. H., An ecological, zoogeographical and taxonomic review of recent planktonic foraminifera. In *Oceanic Micropaleontology* (ed. Ramsay, A. T. S.), Academic Press, London, 1977, vol. 1, pp. 1–100.
15. Prell, W. L. and Curry, W. B., Faunal and isotopic indices of monsoonal upwelling: western Arabian Sea. *Oceanol. Acta*, 1981, **4**, 91–98.
16. Anderson, D. M. and Prell, W. L., A 300 kyr record of upwelling off Oman during the late Quaternary: evidence of the Asian southwest monsoon. *Paleoceanography*, 1993, **8**, 193–208.
17. Overpeck J. T., Anderson, D., Trumbore, S. and Prell, W., The southwest Indian monsoon over the last 18,000 years. *Climate Dyn.*, 1996, **12**, 213–225.
18. Gupta A. K., Anderson, D. M. and Overpeck, J. T., Abrupt changes in the Asian southwest monsoon during the Holocene and their links to the North Atlantic Ocean. *Nature*, 2003, **421**, 354–357.
19. Kroon, D., Distribution of extant planktic foraminiferal assemblages in Red Sea and northern Indian ocean surface waters. *Rev. Española Micropaleontol.*, 1991, **23**, 37–74.
20. Conan, S. M.-H. and Brummer, G.-J. A., Fluxes of planktic foraminifera in response to monsoonal upwelling on the Somalia Basin margin. *Deep-Sea Res. II*, 2000, **47**, 2207–2227.
21. Curry, W. B., Ostermann, D. R., Guptha, M. V. S. and Ittekkot, V., Foraminiferal production and monsoonal upwelling in the Arabian Sea: evidence from sediment traps. In *Upwelling Systems: Evolution since the Early Miocene*, Geological Society Special Publication, US, 1992, vol. 64, pp. 93–106.
22. Bé, A. W. H. and Hutson, W. H., Ecology of planktonic foraminifera and biogeographic patterns of life and fossil assemblages in the Indian Ocean. *Micropaleontology*, 1977, **23**, 369–414.
23. Zheng, F., Li, Q., Li, B., Chen, M., Tu, X., Tian, J. and Jian, Z., A millennial scale planktonic foraminifer record of the mid-Pleistocene climate transition from the northern South China Sea. *Palaeogeogr., Palaeoclimatol., Palaeoecol.*, 2005, **223**, 349–363.
24. Berger, W. H., Sedimentation of planktonic foraminifera. *Mar. Geol.*, 1971, **11**, 325–358.
25. Fairbanks, R. G., Sverdrlove, M., Free, R., Wiebe, P. H. and Bé, A. W. H., Vertical distribution and isotopic fractionation of living planktonic foraminifera from the Panama Basin. *Nature*, 1982, **298**, 841–844.
26. Beveridge, N. A. S. and Shackleton, N. J., Carbon isotopes in recent planktonic foraminifera: A record of anthropogenic CO₂ invasion of the surface ocean. *Earth Planet. Sci. Lett.*, 1994, **126**, 259–273.

27. Bijma, J., Faber, W. W. and Hemleben, Ch., Temperature and salinity limits for growth and survival of some planktonic foraminifers in laboratory cultures. *J. Foramin. Res.*, 1990, **20**, 95–116.
28. Ravelo, A. C., Fairbanks, R. G. and Philander, S. G. H., Reconstructing tropical Atlantic hydrography using planktonic foraminifera and an ocean model. *Paleoceanography*, 1990, **5**, 409–431.
29. Duplessy, J. C., Moyes, J. and Pujol, C., Deep water formation in the North Atlantic Ocean during the last ice age. *Nature*, 1980, **286**, 479–482.
30. Ivanova, E., Schiebel, R., Singh, A. D., Schmiedl, G., Niebler, H. and Hemleben, C., Primary production in the Arabian Sea during the last 135,000 years. *Palaeogeogr., Palaeoclimatol., Palaeoecol.*, 2003, **197**, 61–82.
31. Oberhänsli, H., Bénier, C., Meinecke, G., Schmidt, H., Schneider, R. and Wefer, G., Planktonic foraminifers as tracers of ocean currents in the eastern South Atlantic. *Paleoceanography*, 1992, **7**, 607–632.
32. Emeis, K.-C., Anderson, D. M., Doose, H., Kroon, D. and Schulz-Bull, D., Sea-surface temperatures and the history of monsoon upwelling in the northwest Arabian Sea during the last 500,000 years. *Quaternary Res.*, 1995, **43**, 355–361.
33. Budziak, D., Schneider, R. R., Rostek, F., Muller, P., Bard, E. and Wefer, G., Late Quaternary insolation forcing on total organic carbon and C37 alkenone variations in the Arabian Sea. *Paleoceanography*, 2000, **15**, 307–321.
34. Rostek, F., Ruhland, G., Bassinot, F. C., Beaufort, L., Müller, P. J. and Bard, E., Fluctuations of the Indian monsoon regime during the last 170,000 years: evidence from sea surface temperature, and salinity and organic carbon records. In *Global Precipitations and Climate Change* (eds Besbois, M. and Desalmand, F.), NATO ASI Series 126, Springer, Heidelberg, 1994, pp. 27–51.
35. Rao, K. K., Jayalakshmy, K. V., Kumaran, S., Balasubramanian, T. and Kutti, M. K., Planktonic foraminifera in waters of the Coromandel coast, Bay of Bengal. *Indian J. Mar. Sci.*, 1989, **18**, 1–7.
36. Wiggert, J. D., Jones, B. H., Dickey, T. D., Brink, K. H., Weller, R. A., Marra, J. and Codispoti, L. A., The northeast monsoon's impact on mixing, phytoplankton biomass and nutrient cycling in the Arabian Sea. *Deep-Sea Res. II*, 2000, **47**, 1353–1385.
37. Weller, R. A. *et al.*, Moored observations of upper-ocean response to the monsoons in the Arabian Sea during 1994–1995. *Deep-Sea Res. II*, 2000, **49**, 2195–2230.
38. Veldhuis, M. J. W., Kraay, G. W., Van Bleijswijk, J. D. L. and Baars, M. A., Seasonal and spatial variability in phytoplankton biomass, productivity and growth in the northwestern Indian Ocean: the southwest and northeast monsoon, 1992–1993. *Deep-Sea Res. I*, 1997, **44**, 425–449.
39. Hansell, D. A. and Peltzer, E. T., Spatial and temporal variations of total organic carbon in the Arabian Sea. *Deep-Sea Res. II*, 1998, **45**, 2171–2193.
40. Dickey, T. *et al.*, Seasonal variability of bio-optical and physical properties in the Arabian Sea: October 1994–October 1995. *Deep-Sea Res. II*, 1998, **44**, 2001–2025.
41. Schott, F. A., Xie, S.-P. and McCreary Jr, J. P., Indian Ocean circulation and climate variability. *Rev. Geophys.*, 2009, **47**, RG1002.

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Emerging new multi-drug resistant bacterial pathogen, *Acinetobacter baumannii* associated with snakehead *Channa striatus* eye infection

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Although *Acinetobacter baumannii* acts as a severe human pathogen, there are only few studies to date that report it as a pathogen for fish. In the present study, one virulent bacterial strain was isolated from diseased *Channa striatus*, from a farm at the Central Institute of Freshwater Aquaculture, Bhubaneswar, Orissa, which showed symptoms like cloudy eyes, pop eye (exophthalmia), opaque lenses and mild ulceration on the whole body irrespective of sex and size of fish. Based on morphology, different biochemical tests and sequence analysis of 16S rDNA segment as well as phylogenetic study, the causative bacterium (called ChE) was identified as *A. baumannii*. The pathogenicity was further confirmed by artificial infectivity study (LD₅₀ dose of 10^{8.37} CFU/fish). In the drug sensitivity study, this isolate was highly resistant to many antibiotics. The isolate was also highly resistant to all three tested heavy metals (Cu⁺², Cr⁺⁶, Hg⁺²), thus proving its virulent nature.

Keywords: *Acinetobacter baumannii*, *Channa striatus*, drug resistant, pathogen.

THE genus *Acinetobacter* belonging to the family Moraxellaceae, within the gamma subdivision of proteobacteria is ubiquitous in nature as it is found frequently in soil, water and dry environments^{1–5}. It is a group of Gram-negative, strictly aerobic, non-motile coccobacilli. Currently, *Acinetobacter* sp. is emerging as a serious human nosocomial pathogen being involved in several infections, e.g. bacteremia, urinary tract infection, secondary meningitis and ventilator-associated pneumonia². This species, especially *Acinetobacter baumannii* is treated as the most clinically important microorganism due to its remarkable ability to develop resistance to many antibiotics^{6–11}. There are some reports on the incidence of infection in fish, which suggest that it could also be treated as severe fish pathogen. The most probable first report was from China and the bacterial strain isolated from mandarin fish (*Siniperca chutasi*) was confirmed as *A. baumannii* in terms of its biochemical characteristics¹². Another report was also from China and the species isolated from diseased channel catfish proved as the virulent pathogen for this fish¹³. In the earlier reports, the methods of identification

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